



spatiotemporal distribution of these events directly reflects the propagation and complexity of HF-induced fractures. Passive microseismic monitoring (PMM) thus serves as a critical tool for evaluating HF effectiveness, enabling the characterization of fracture geometry, fluid migration pathways, and stimulated reservoir volume (SRV) (Maxwell and Urbancic, 2001). Downhole PMM captures signals by arranging multiple receivers in one or multiple monitoring wells adjacent to the treatment well, which is considered to be the most reliable monitoring method due to the proximity of the monitoring system to seismic source, low environmental noise, and high signal-to-noise ratio (SNR). Notably, while various unconventional reservoirs exhibit distinct seismic

## 2. Methodology

Microseismic data processing exhibits inherent nonlinearity characteristics and high uncertainty, convolutional data-driven paradigm based on point estimation is difficult to effectively capture uncertainty information, consequently diminishing interpretability of prediction results. To address above limitations, it is of great value to systematically carry out the uncertainty source analysis of the DL-based method and realize the uncertainty quantification. In this section, we first construct a data-driven framework integrating uncertainty quantification strategy for phase detection and arrival picking, and then form a hybrid-driven workflow for comprehensive microseismic data processing.

### 2.1. Uncertainty-aware DL framework for phase detection and arrival picking

Microseismic events induced by HF are usually characterized by low energy magnitude and high waveform similarity with ambient noise, consequently enhancing detection sensitivity becomes paramount, particularly for low SNR events. Currently, in the field of artificial intelligence (AI) seismology, the application of DL for seismic phase detection and arrival picking has gained wide academic and industrial recognition compared to other application scenarios. This comparative advantage originates from two key factors: (1) the inherent similarity between seismic phase detection tasks and computer vision-based pattern recognition, enabling straightforward migration of mature algorithms that significantly reduce the technical barriers of model construction; and (2) the availability of substantial labeled samples that facilitate training supervised machine learning models with strong generalization abilities. Meanwhile, regarding downhole microseismic monitoring systems, research focusing on DL-based model constructions and associated scientific issues has emerged as a critical direction.

#### 2.1.1. Monte Carlo Dropout strategy

Uncertainty primarily encompasses two types in DL (Fig. 1): epistemic uncertainty and aleatoric uncertainty (Kiureghian and Ditlevsen, 2009). The former stems from the model's limited knowledge and can be mitigated by enhancing the model architecture or expanding training data, while the latter arises from inherent noises or ambiguity in the data, requiring explicit incorporation of uncertainty quantification during model design. The distinction between these two uncertainties in a regression framework is illustrated in Fig. 1(a). The black dashed curve represents the ground-truth function  $y = f(x)$ . Blue error bars denote observed data points contaminated with Gaussian noises, characterizing aleatoric uncertainty inherent to measurement processes. The gray shaded area delineates the 2- $\sigma$  confidence interval of a model's predictions, reflecting epistemic uncertainty due to incomplete knowledge. The uncertainty peaks near  $x = 5$ , simulating a localized knowledge gap (e.g., sparse training data or model inadequacy). Critically, aleatoric uncertainty persists globally as irreducible noises, while epistemic uncertainty diminishes with improved models or additional data. Consequently, our research focuses on quantifying the latter through systematic analysis.

Dropout is a widely adopted regularization technique in DL, first proposed by Srivastava et al. (2014) as a means to mitigate overfitting in neural networks. The core principle involves the stochastic deactivation of neurons during the training phase, where each neuron in a fully connected layer is temporarily

dropped with a predefined probability  $p$ . This process introduces randomness that prevents the DL model from becoming overly dependent on specific neurons or local features, thereby improving its generalization performance.

Mathematically, for a neural network layer  $l$  (totally  $L$ ) with 2D weight matrix  $\mathbf{W}^{(l)} \in \mathbf{C}^{m \times n}$ , input  $\mathbf{H}^{(l-1)}$  and bias vector  $\mathbf{b}^{(l)}$ , the conventional output  $\mathbf{H}^{(l)}$  can be denoted as:

$$\mathbf{H}^{(l)} = \left( \mathbf{W}^{(l)} \mathbf{H}^{(l-1)} + \mathbf{b}^{(l)} \right) \quad (1)$$

where  $\sigma$  is the activation function. Then, the above form of dropout is modified by element-wise multiplication with a Bernoulli-distributed mask  $\Phi^{(l)}$ :

$$\mathbf{H}_D^{(l)} = \left( \mathbf{W}_D^{(l)} \mathbf{H}_D^{(l-1)} + \mathbf{b}^{(l)} \right)$$

$$\mathbf{W}_D^{(l)} = \mathbf{W}^{(l)} \odot \Phi^{(l)} = \begin{bmatrix} w_{11} & w_{12} & \dots & w_{1n} \\ w_{21} & w_{22} & \dots & w_{2n} \\ \dots & \dots & \dots & \dots \\ w_{m1} & w_{m2} & \dots & w_{mn} \end{bmatrix}$$

$$\Phi^{(l)} \sim \text{Bernoulli}(1 - p) \quad (2)$$

Here,  $\odot$  denotes the Hadamard product (element-wise multiplication), where each entry of the resulting matrix is the product of the corresponding entries in  $\mathbf{W}^{(l)}$  and  $\Phi^{(l)}$ . During training, the mask is regenerated for each mini-batch, creating an ensemble effect where the network effectively learns multiple subnetworks simultaneously. The theoretical foundation of dropout can be interpreted through two complementary perspectives:

- Co-adaptation prevention: By disrupting fixed interaction patterns between neurons, dropout forces individual units to develop robust features independently of specific collaborators.
- Implicit model averaging: The training process approximates the optimization of an ensemble of multiple subnetworks, effectively performing approximate Bayesian marginalization over network structures.

Monte Carlo Dropout is an extension of the traditional dropout technique, proposed by Gal and Ghahramani (2016). Its core idea lies in extending the stochastic deactivation mechanism from the training phase to the prediction phase, enabling approximate Bayesian inference through multiple forward propagation samplings to quantify prediction uncertainty. Formally, this corresponds to learning a variational distribution that minimizes the KL divergence from the true parameter posterior. Fig. 1(b) illustrates the difference between dropout and MC-Dropout strategies during the training and prediction phases. Unlike dropout, which is only active during training, MC-Dropout retains stochastic neuron deactivation during prediction, forming a Monte Carlo estimation of the posterior distribution.

Let the trained neural network parameters (weight matrix  $\bar{\mathbf{W}}$  and bias  $\bar{\mathbf{b}}$ ) operate on the input  $x$ . In traditional inference, the output  $\mathbf{H}^{(l)}$  is computed via deterministic forward propagation. In contrast, MC-Dropout generates independent Bernoulli masks during each forward pass and repeats sampling  $T$  times. The predictive distribution is denoted as:

Fig. 1. Uncertainty characterization in DL: (a) aleatoric vs epistemic uncertainties and (b) dropout/MC-Dropout strategies during the training and prediction phases.

$$\mathbf{H}_{MD}^{(L)}(t) = \frac{(\mathbf{H}_{MD}|x, \overline{\mathbf{W}}(t), \overline{\mathbf{b}}(t))}{\overline{\mathbf{W}}(t) = \overline{\mathbf{W}} \odot \boldsymbol{\varepsilon}(t)} \quad (3)$$

where  $\overline{\mathbf{W}}(t)$  and  $\overline{\mathbf{b}}(t)$  represent the stochastic weights and bias for the  $t$ -th sampling, respectively. The mean (Me) and standard deviation (Sd) of the output  $\mathbf{H}_{MD}$  can be calculated as:

$$\begin{aligned} \text{Me}(\mathbf{H}_{MD}) &\approx \frac{1}{T} \sum_{t=1}^T \mathbf{H}_{MD}^{(L)}(t), \\ \text{Sd}(\mathbf{H}_{MD}) &\approx \sqrt{\frac{1}{T} \sum_{t=1}^T \left( \mathbf{H}_{MD}^{(L)}(t) - \text{Me}(\mathbf{H}_{MD}) \right)^2} \end{aligned} \quad (4)$$

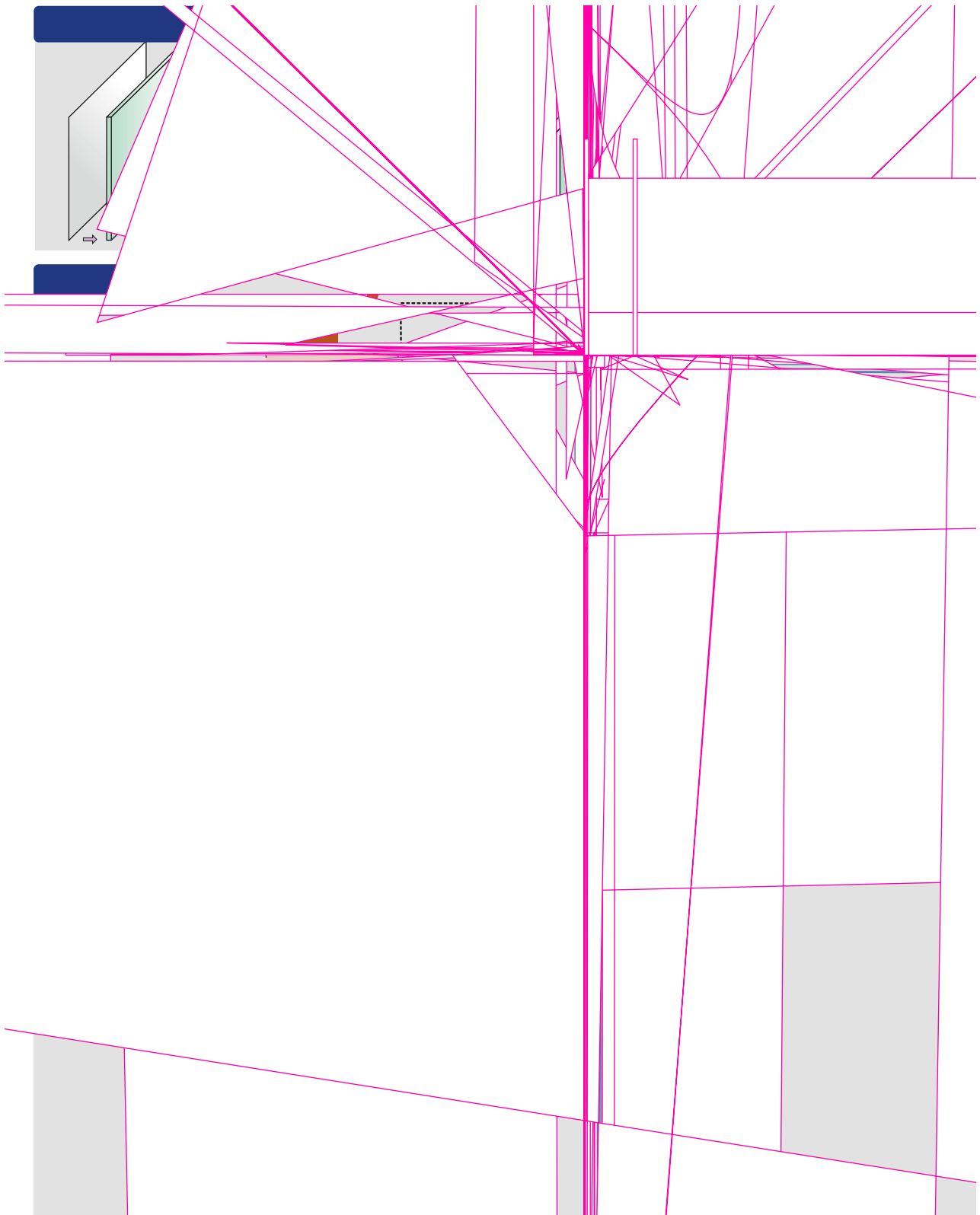
### 2.1.2. MD-Net

The U-Net model, a convolutional neural network architecture, was initially proposed by [Ronneberger et al. \(2015\)](#) for biomedical image segmentation tasks, demonstrating remarkable sensitivity in detecting edge features. Analogous to image segmentation, the phase detection and arrival picking of P-waves and S-waves can be conceptualized as a multi-class classification optimization problem. Within this framework, each sampling point in seismic waveforms should be categorized into one of three classes: P-wave, S-wave, or noise. Our study aims to achieve accurate classification of all sampling points while detecting edge characteristics at the boundaries between noise and valid seismic phases effectively, thereby enabling precise determination of phase arrival times. Based on the U-Net, we propose a new framework named MD-Net (Monte Carlo Dropout-Net) which constructs a feature-enhanced model using multi-trace microseismic recordings. By inputting target multi-trace waveform data into this

established model, the workflow of phase detection and arrival picking can be finished automatically.

The network architecture of MD-Net is illustrated in [Fig. 2\(a\)](#). The model takes multi-trace segmented waveforms extracted from continuous microseismic recordings as input units, where three channels correspond to two horizontal components (E-W and N-S) and one vertical component (Z) of waveforms. The output layer generates probability distribution maps with identical dimensions to the input layer, comprising three independent channels representing probability distributions for P-wave, S-wave, and noise, respectively. Structurally, both input and output data are organized as 3D tensors of size  $N \times M \times K$ . The first dimension ( $N$ ) denotes the “Intra-Trace” dimension, specifying the number of sampling points per trace, which should theoretically ensure complete inclusion of double-phase seismic events (e.g., full waveforms of both P- and S-phases). The second dimension ( $M$ ) represents the “Inter-Trace” dimension, indicating the number of traces for joint processing, with its value determined by the actual spatial arrangement of receivers in the observation system. The third dimension ( $K$ ) corresponds to the “Channel” dimension, representing three-component seismic recordings (two horizontal components and one vertical component), hence maintaining a fixed value of 3.

The MD-Net architecture is composed of symmetrical encoding and decoding paths, implementing hierarchical feature extraction and reconstruction through multiple downsampling and upsampling stages. The encoding path consists of five cascaded downsampling stages, each containing a convolutional layer and a pooling layer: (a) The 2D convolutional layer employs multiple convolutional kernels to extract spatiotemporal features. Convolution operations along the first dimension (temporal axis) and second dimension (spatial axis) jointly model “Intra-Trace” waveform characteristics and “Inter-Trace” phase arrival



**Fig. 2.** MD-Net framework for uncertainty-aware phase detection and arrival picking: (a) network architecture; (b) training phase integrating Monte Carlo Dropout sampling and multi-loss optimization; and (c) prediction phase with uncertainty quantification (mean value [Me] and confidence interval [Me–Sd, Me + Sd] derived from  $T$  stochastic outputs).

correlations. (b) The max-pooling layer performs spatial downsampling on feature maps, preserving salient features while strategically reducing spatial resolution to control parameter growth. All convolutional kernels are configured with a size of (10, 2). The pooling factor is set to (5, 1) for the first two pooling layers and (2, 1) for the subsequent three layers. Through the decoding path, the single-channel sampling resolution of input data is progressively reduced to 1/200 of the initial value ( $N/200$ ). To compensate for information loss, the network implements a channel expansion strategy synchronously to increase feature channels from 3 to 128. Each upsampling stage in the decoding path contains a transposed convolutional layer and a feature fusion layer: (a) Spatial dimensions are restored by the transposed convolution operation through deconvolutional kernels, with magnification factors strictly equivalent to the corresponding pooling factors in the encoding path. (b) The skip connection operation for feature fusion establishes cross-layer feature reuse by concatenating the same-scale feature maps from the downsampling and upsampling stages, ultimately constructing an end-to-end encoder-decoder architecture analogous to the U-Net framework. The deconvolutional kernels maintains parameter symmetry with convolutional kernels (kernel size = (10, 2)) to progressively recover the feature map to the original spatial resolution ( $N \times M \times K$ ). All convolutional and deconvolutional convolutional layers employ ReLU nonlinear activation functions (Glorot et al., 2011). Finally, a Softmax layer converts the output into a probability distribution within the interval [0, 1]. The specific functional form adopted in this study is:

$$q_i(\text{label}) = \text{Softmax}(Z) = \frac{e^{Z_i(\text{label})}}{\sum_{j=1}^3 e^{Z_j}} \quad (5)$$

where  $Z_i$  denotes the input value at the  $i$ -th sampling point of the Softmax layer. The labels 1–3 correspond to the P-wave, S-wave, and noise channels respectively, with  $q$  representing the output probability distribution among these three channels.

Consistent with the output probability distribution  $q$ , the sample labels also adopt probability distributions for the P-wave, S-wave, and noise channels. For samples including P- or S-wave arrivals requiring simultaneous phase detection and arrival picking, the three-channel probability distributions  $p(P)$ ,  $p(S)$ , and  $p(N)$  are defined using manual phase arrival pickings as reference standards. Taking a typical single-trace double-phase sample label (Fig. 2(b)) as an example, we account for potential inaccuracies in manual phase arrival pickings (represented by red vertical lines in the figure: solid for P-wave arrival, dashed for S-wave arrival). The P-wave and S-wave labels are designed as Gaussian probability distributions centered at their respective arrivals ( $\mu(P)$  for P-wave,  $\mu(S)$  for S-wave), with standard deviations  $\sigma(P)$  and  $\sigma(S)$ . The formulas for  $p(P)$ ,  $p(S)$ , and  $p(N)$  at the  $i$ -th sampling point are given by:

$$\begin{cases} p_i(P) = \exp\left(-\frac{(i - \mu(P))^2}{2(\sigma(P))^2}\right) \\ p_i(S) = \exp\left(-\frac{(i - \mu(S))^2}{2(\sigma(S))^2}\right) \\ p_i(N) = 1 - p_i'(P) - p_i'(S) \end{cases} \quad (6)$$

For noise samples, they just need to be detected correctly. Thus, we set  $p(P) = 0$ ,  $p(S) = 0$ , and  $p(N) = 1$ .

To integrate the Monte Carlo Dropout strategy into the network for uncertainty estimation, we introduce five dropout layers after each pooling operation in all downsampling stages and after skip

connection operations in all upsampling stages. During both training and prediction phases of the MD-Net model, the dropout layers must remain activated (Gal and Ghahramani, 2016). The dropout rate directly influences the sensitivity of uncertainty estimation. If set too high, the model may become overly conservative, potentially obscuring genuine seismic phase characteristics (e.g., weak P-wave signals). Conversely, if set too low, the model may fail to adequately capture uncertainty caused by noise interference, thereby reducing robustness for low SNR events. Generally, low-SNR samples require higher dropout rates. In practice, differentiated dropout rates should be implemented based on model complexity and dataset qualities with varying SNR levels.

During training of the MD-Net model, a loss function is defined to quantify the discrepancy between model outputs and sample labels, with the optimization objective of minimizing this function. For the multi-class classification task in this study (identifying P-wave, S-wave, and non-phase noise in seismic waveforms), the standard categorical cross-entropy loss combined with a Softmax activation function can measure the discrepancy between the predicted probability distribution and true labels, expressed as:

$$L_{p,q} = - \sum_{\text{label}} \sum_i p_i(\text{label}) \log q_i(\text{label}) \quad (7)$$

However, this task suffers from a severe class imbalance issue: non-phase noise segments dominate the waveform data (e.g., in double-phase events shown in Fig. 2(b), P- and S-wave phases constitute ~10%, while noise accounts for ~90%). To address this, we propose a weighted cross-entropy loss function that dynamically assigns distinct weighting factors to phase and non-phase segments, thereby increasing the penalty for errors in phase segments to enhance the loss function's sensitivity to phase-related discrepancies (Fig. 2(b)). The loss function  $L$  is formulated as:

$$L_{p,q} = L_{p,q}^{(P)} + L_{p,q}^{(S)} + L_{p,q}^{(N)} \quad (8)$$

where the sub-terms are defined as:

$$\begin{cases} L_{p,q}^{(P)} = - \sum_i [ \alpha_0 p_i(P) \log q_i(P) + \alpha_1 (1 - p_i(P)) \log(1 - q_i(P)) ] \\ L_{p,q}^{(S)} = - \sum_i [ \alpha_0 p_i(S) \log q_i(S) + \alpha_1 (1 - p_i(S)) \log(1 - q_i(S)) ] \\ L_{p,q}^{(N)} = - \sum_i [ \alpha_1 p_i(N) \log q_i(N) + \alpha_0 (1 - p_i(N)) \log(1 - q_i(N)) ] \end{cases} \quad (9)$$

Here,  $\alpha_0$  and  $\alpha_1$  are dynamic weighting factors with  $\alpha_0 > \alpha_1$ , ensuring that prediction errors for phase segments (P- and S-waves) contribute more significantly to the total loss.

For a microseismic input sample, performing  $T$  forward propagation processes during prediction phase yields  $T$  output volumes with dimensions  $N \times M \times 3$  (Fig. 2(c)). The final probability distribution is derived by computing Me and Sd across these outputs. Key points of MD-Net include:

- (a) Phase detection: For one microseismic channel of a sample, we record the maximum value of the mean distribution and its corresponding sampling point position. For a sample, we compare the maximum values of the P-wave and S-wave mean probability distributions against preset thresholds to classify events as double-phase, single-phase, or noise. The setting of thresholds for P-wave and S-wave detection is critical. Excessively high thresholds may result in the misclassification of low SNR seismic phases as noise, while overly low thresholds could erroneously detect noise events

as valid seismic signals. The specific values for thresholds must be determined through case-by-case analysis, tailored to the specific characteristics of datasets and scenarios.

- (b) Phase arrival picking: For one double-phase (including P-wave and S-wave) or single-phase event (including only P-wave or S-wave), the position corresponding to the maximum value in the mean probability distribution is defined as the predicted phase arrival time. The theoretical foundation lies in the fact that the statistical average of multiple sampling results can effectively suppress random noise interference, causing the position of the probability peak to converge towards the true arrival time.
- (c) Uncertainty awareness: The interval  $U = [Me - Sd, Me + Sd]$  provides a quantitative estimate of prediction uncertainty for each valid seismic phase. The range of  $U$  reflects the magnitude of uncertainty for the prediction results. By overlaying  $U$  onto microseismic waveforms, we can distinguish among “high-uncertainty phase arrivals” (likely caused by noise interference) and “low-uncertainty phase arrivals” (enough reliable seismic phases). By utilizing the acquired uncertainty information, we can improve the accuracy of phase arrival picking through manual calibration of model prediction results. Furthermore, applying the calibrated results to subsequent inversion processes contributes to enhancing the reliability of source location.

## 2.2. Workflow optimization for source location

The traditional model-driven paradigm suffer from inherent limitations due to their heavy dependence on a priori knowledge and oversimplification of complex geological conditions. In contrast, the data-driven paradigm significantly improve performance in complex scenarios by learning implicit patterns from large-scale datasets. Nevertheless, traditional paradigm can still be useful in interpretable and small-sample scenarios in combination with data-driven paradigm. The comparison of different paradigms is summarized in [Table 1](#).

In PMM, rapid phase detection, arrival picking, and event location are needed and its core objective is to establish an automated processing workflow that is efficient, accurate, and has minimal

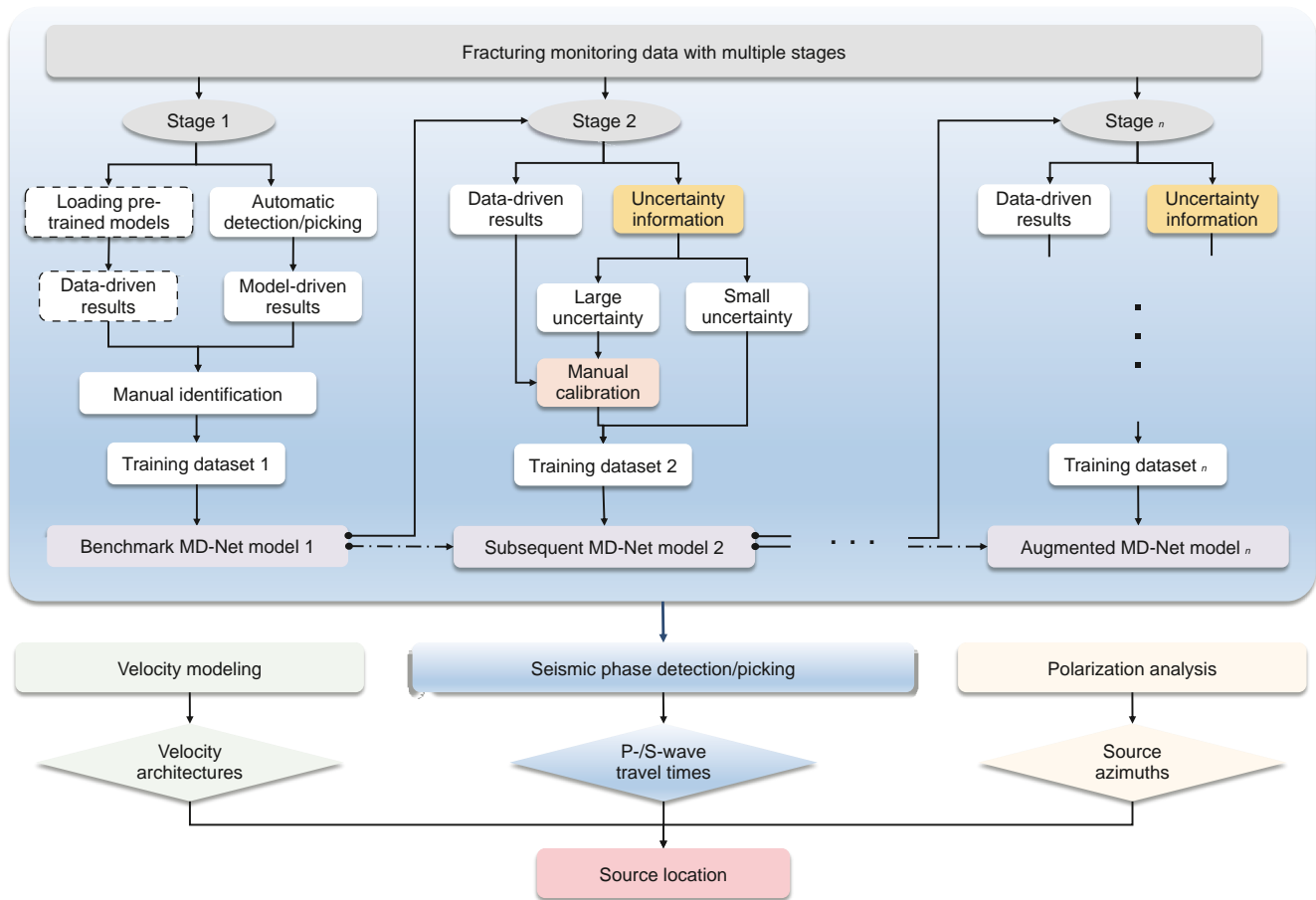


Fig. 3. Workflow optimization for source location of multi-stage fracturing monitoring using the hybrid driven paradigm.

function. The optimal source location is subsequently determined through iterative optimization search within the 3D space. For specifics on source azimuth determination, objective function formulation, and the search algorithm utilized, see [Tan et al. \(2017\)](#). The advantages of this hybrid-driven workflow are: (1) Utilizing the law of data acquisition in real-time data processing to construct an adaptive and automated process. The model of phase detection and arrival picking are constructed progressively during HF, and the model parameters are modified and the generalization ability is continuously enhanced. (2) The synergy between manual quality control and automated methods improves the accuracy while significantly reducing the uncertainty. (3) The innovative combination of model-driven and data-driven paradigms realizes the complementary advantages and forms a domain-adaptable hybrid-driven optimization framework.

In the workflow, manual quality control is systematically designed to minimize subjective judgment through standardized procedures. For samples predicted by MD-Net as high-uncertainty (confidence intervals exceeding predefined thresholds), independent manual annotations are performed by multiple geophysical analysts. These analysts adhere to unified guidelines for phase detection and arrival picking, ensuring consistency in labeling criteria. Their individual annotations are subsequently aggregated through consensus-based reconciliation. This multi-operator cross-validation approach effectively mitigates individual biases, reduces arbitrariness in ambiguous cases, and enhances the

reproducibility of processing results. The finalized annotations are then fed back to update model, establishing a closed-loop workflow.

### 3. Field applications

#### 3.1. Dataset and preprocessing

The dataset is acquired from a PMM program during one horizontal well HF project in a tight oil and gas field in eastern China, containing continuous monitoring recordings of 10 fracturing stages over approximately 16 h. The observation system consists of 15 three-component receivers deployed in an adjacent vertical monitoring well with 10 m spacing, operating at the sampling rate of 2000 Hz. [Fig. 4\(a\)](#) shows the observation system where target fracturing zones are located at about 2700 m depth, while the receiver array covers the interval of 2443–2673 m, with the location of ball-drop events indicated by red asterisks.

The preprocessing for raw microseismic recordings comprises three steps: (1) Spectral filtering: Application of a 30–350 Hz band-pass filter to suppress low-frequency ambient noise and high-frequency instrumental artifacts. (2) Receiver orientation: Alignment of horizontal components using perforation events ([Fig. 4\(b\)](#)), resulting in enhanced waveform coherence and sharper phase arrivals. (3) Data standardization: Segmentation using a 1200-sample window (0.6 s) based on the observed P-S arrival time differences (<0.25 s), followed by demeaning and amplitude normalization via standard deviation.

The MD-Net architecture features an encoder-decoder structure with the input dimension of  $1200 \times 15 \times 3$ . Firstly, the number of sampling points  $N$  within a single channel is reduced from 1200 to 6, and the number of channels  $K$  is increased from 3 to 128 by 5 downsampling stages; and then the original dimension is recovered gradually by 5 upsampling stages. The weighting factors  $w_0 = 6.7$  and  $w_1 = 1.2$  are derived from the inverse proportion of seismic phase segments (P-/S-waves, ~15%) to noise segments (~85%) in one waveform sample, prioritizing penalties for seismic phase segments.

Four sample categories are defined: background noise, noise event, single-phase event, and double-phase event. Following monitoring practice, only single- and double-phase events are retained as valid samples, with others classified as noises in the dataset construction. Manual arrival picking uncertainties were set at  $\tau_p = 10$  ms (20 samples) and  $\tau_s = 20$  ms (40 samples) respectively, reflecting the superior P-wave detectability due to higher SNR level (Fig. 2(b)).

The majority of the segmented recordings consist of noises, with valid seismic phase samples representing only a small fraction. If all samples from the continuous recordings are used as inputs for model training, a typical sample imbalance problem in DL may occur, resulting in poor performance for the valid seismic phase samples (López et al., 2013). To address this issue, a sampling strategy was employed to select a subset of noise samples alongside all available seismic phase samples for the training dataset. Ultimately, it was determined that the training dataset originated from the first seven fracturing sta [A-2]Z\_(S(

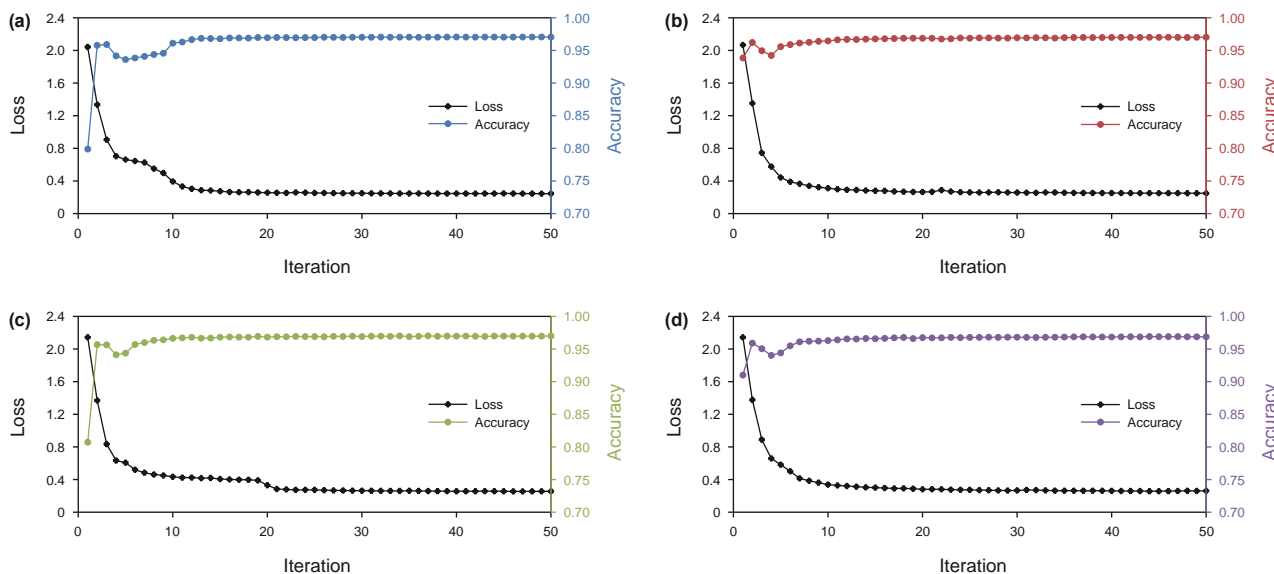


Fig. 5. Loss and accuracy curves during training: (a) dropout rate = 0, (b) dropout rate = 0.1, (c) dropout rate = 0.2, and (d) dropout rate = 0.3.

Table 2 Architectural differences among experimental groups.

Group	Training phase	Prediction phase	Forward propagation samplings	Uncertainty quantification
T-Base	No dropout	No dropout	Deterministic (single time)	Not applicable
T-Dr series	Dropout (rate: 0.1/0.2/0.3)	No dropout	Deterministic (single time)	Not applicable
T-MD series	Dropout (rate: 0.1/0.2/0.3)	Dropout (rate: 0.1/0.2/0.3)	Stochastic (multiple times)	Enabled via ensemble mean

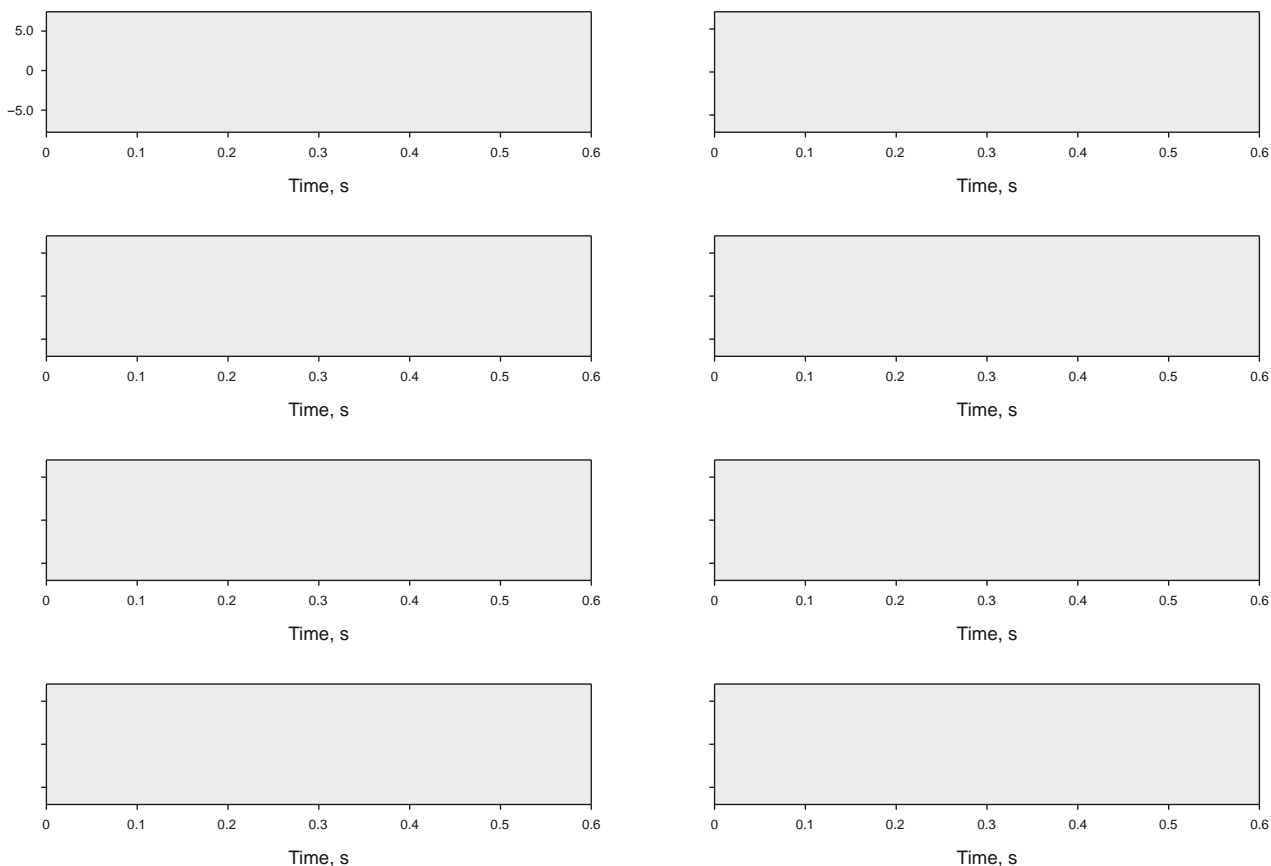
wave arrival (Fig. 6(b)). For phase arrival picking, T-Base mistakenly picks the P-wave to a wrong position, but the picking results of the rest of groups do not notably deviate from the true arrival. Further analysis shows that there is a typical noise interference near the position of P-wave arrival picked by T-Base, and the jumping phenomenon can also be observed there in Fig. 6(c), (d), (f), and (h). The uncertainty estimation (Fig. 7) reveals elevated uncertainties near phase arrivals, particularly for the P-wave, mostly attributable to the typical noise interference mentioned above. This demonstrates MC-Dropout’s double advantages: mitigating large picking errors through stochastic regularization while providing quantitative reliability metrics for prediction results, which is crucial for the subsequent manual calibration.

The second sample presents a low SNR trace (Figs. 8 and 9), where the P-wave arrival is entirely obscured by background noise (Fig. 8(a)). This challenging case demonstrates phase detection capability even for manual analysis, requiring integration with the similarity of adjacent trace waveforms or template matching techniques for reliable detection. As shown in Fig. 8(b)–(h), all groups exhibit significant misclassification manifested through phase confusion between P- and S-waves, irregular output probability distribution morphologies, and substantial deviation of arrival time picks. The T-MD series display elevated prediction uncertainties throughout the seismic phases (Fig. 9(b)–(d)). Detailed analysis of the true P-wave arrival position reveals distinct behavioral patterns: The T-MD1 shows no significant jumping in either mean probability distribution (Fig. 8(d)) or uncertainty estimation (Fig. 9(b)), indicating background noise. In contrast, T-MD2 and T-MD3 demonstrate measurable probability jumping (Fig. 8(f) and (h)) accompanied by elevated uncertainties (Fig. 9(c) and (d)), indicating uncertain phase. This case highlights the necessity for synthesizing uncertainty metrics from multiple

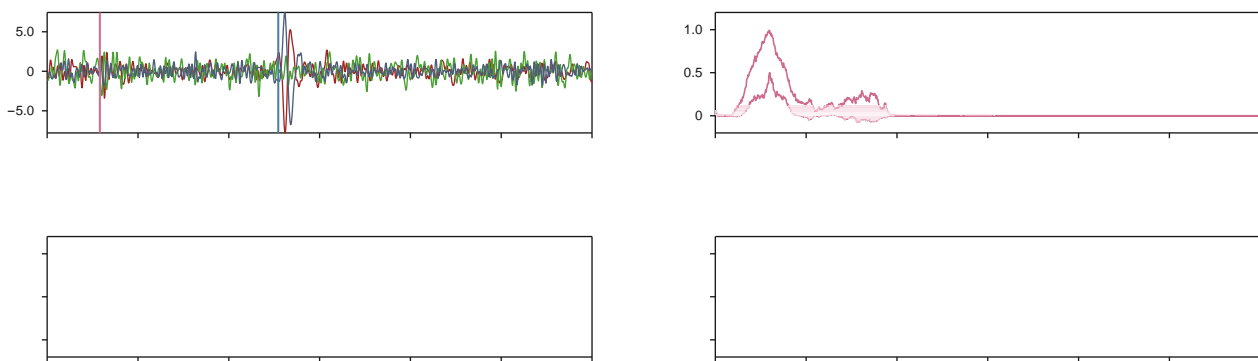
MD-Net configurations with different dropout rates and the waveform characteristics among traces for manual analysis.

The third sample presents another low SNR trace (Figs. 10 and 11). The P-wave phase could be correctly detected by all experimental groups, while the T-Base, T-Dr1, and T-MD1 do not have a significant jump near the S-wave arrival, erroneously classifying this sample as a single-phase event. Groups with higher dropout rates (T-Dr2/3, T-MD2/3) maintain better detection performance near the S-wave arrival, consistent with the generalization enhancement effect of stronger stochastic regularization. In picking accuracy, T-Dr2 and T-MD2 achieve precise S-wave picking, whereas T-Dr3 and T-MD3 exhibit insufficient probability jumping at true S-wave arrival with notable picking deviations. This sensitivity to dropout rate hyperparameter underscores the importance of maintaining appropriate regularization intensity during model configuration. Uncertainty visualization (Fig. 11(b)–(d)) confirms the essential role of manual calibration for such ambiguous low SNR cases, with prediction reliability metrics providing crucial decision-support information for practical scenarios.

From results, all experimental groups achieved zero false positives in detecting noise samples as valid microseismic events (including single-phase and double-phase events) across the entire testing dataset, demonstrating MD-Net’s robustness against false alarms under diverse noise conditions. Systematic evaluation of microseismic samples across Figs. 6–11 reveals two critical findings: Firstly, the MC-Dropout technique enables effective uncertainty quantification without compromising phase detection or arrival picking accuracy. Secondly, the relationship between dropout rate and model generalization follows a nonlinear pattern, necessitating systematic optimization through cross-validation procedures.



**Fig. 6.** Phase detection and arrival picking results of different groups (double-phase example 1): (a) three component single-trace seismogram, (b) T-Base, (c) T-Dr1, (d) T-MD1, (e) T-Dr2, (f) T-MD2, (g) T-Dr3, and (h) T-MD3. In (b), (c), (e), and (g), the red and blue curves represent the output probability distributions of P- and S-waves, respectively; while in (d), (f), and (h), the curves represent the Me values of the output probability distributions over  $T$  iterations. All subfigures show manual picks (solid vertical lines) and model-predicted picks (dashed vertical lines), with probability thresholds marked by horizontal black dashed lines.



**Fig. 7.** Uncertainty quantification of T-MD series (double-phase example 1): (a) three component single-trace seismogram, (b) T-MD1, (c) T-MD2, and (d) T-MD3. In (b), (c), and (d), intervals  $[Me-Sd, Me+Sd]$  between red/blue curves represent uncertainty bounds of P-/S-waves.

The detection accuracy metric is defined as the ratio of correctly detected traces to total double-phase traces of the testing dataset. The detection performance of double-phase samples is summarized in Table 3. All experimental groups exhibited detection omissions, with T-Base demonstrating the poorest performance at only 1244 correctly detected traces. The T-Dr and T-MD series show progressive accuracy improvements. With a detection accuracy of 99.00% as the threshold, the high-accuracy group

(>99.00%) consists of four configurations with dropout rates of 0.2 and 0.3 (T-Dr2/3, T-MD2/3), while the low-accuracy group (<99.00%) consists of three configurations with dropout rates of 0 and 0.1 (T-Base, T-Dr1, and T-MD1). This statistical analysis demonstrates that the dropout rate serves as a critical controlling factor for model detection performance.

Fig. 12 displays phase arrival picking error distributions relative to manual benchmarks, with quantitative metrics detailed in

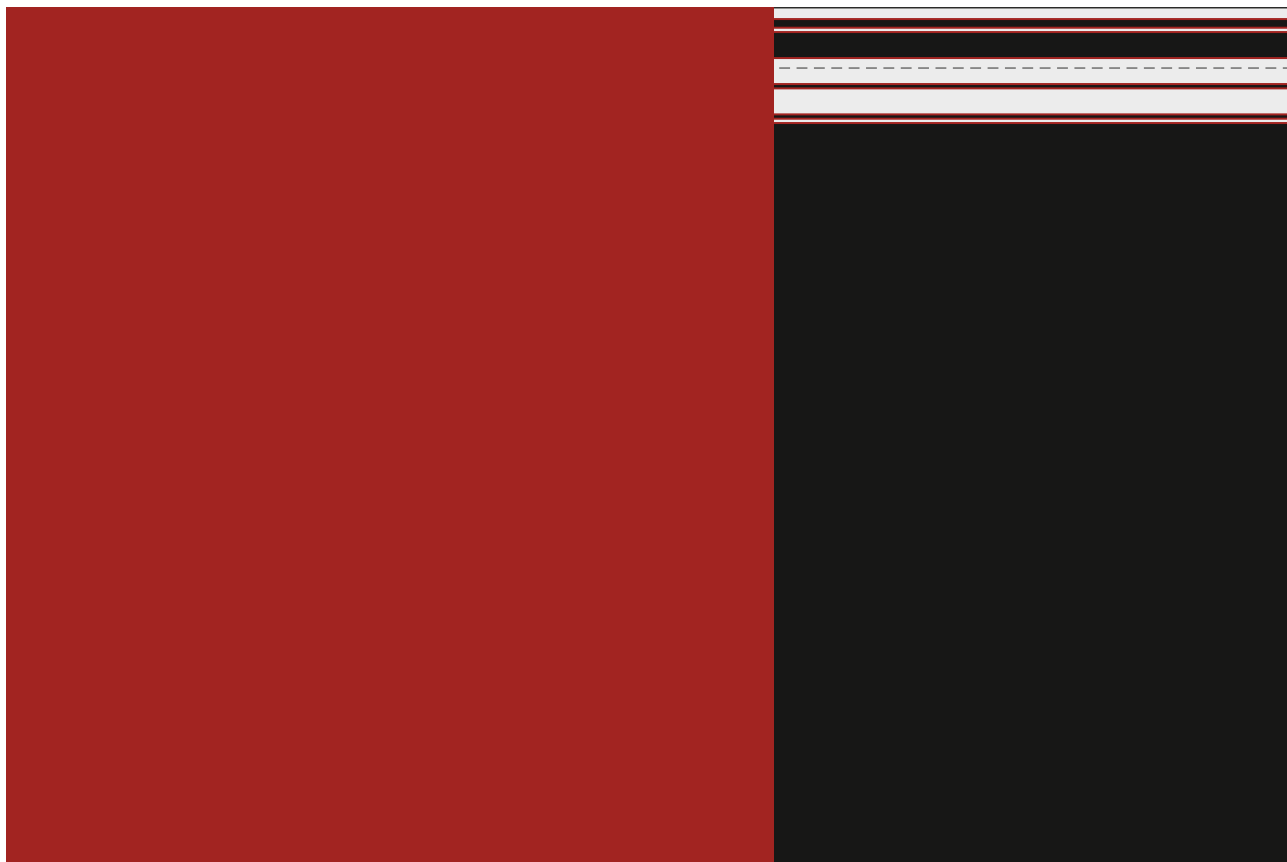


Fig. 8. Phase detection and arrival picking results of different groups (double-phase example 2): (a) three component single-trace seismogram, (b) T-Base, (c) T-Dr1, (d) T-MD1, (e) T-Dr2, (f) T-MD2, (g) T-Dr3, and (h) T-MD3.

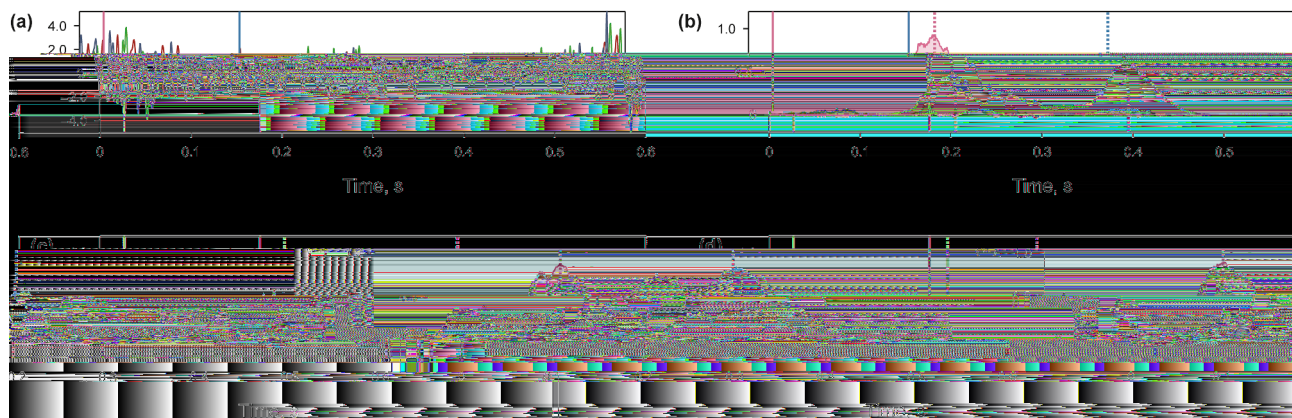
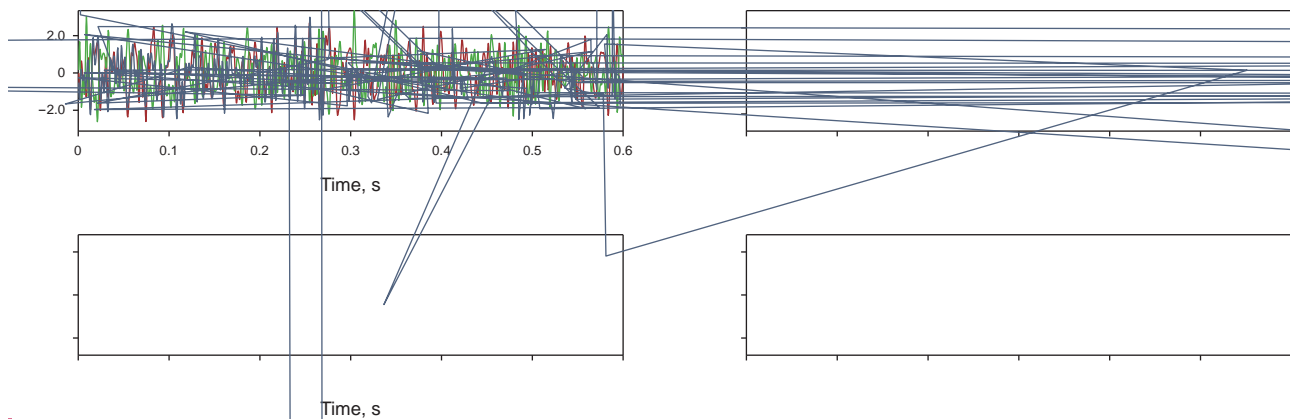


Fig. 9. Uncertainty quantification of T-MD series (double-phase example 2): (a) three component single-trace seismogram, (b) T-MD1, (c) T-MD2, and (d) T-MD3.

Table 4. All groups achieve the mean of picking errors (ME) within 0–1.7 ms for P-wave, S-wave, and P-S time difference, satisfying microseismic monitoring precision requirements. Standard deviation of picking errors (SD) analysis demonstrates significant error reduction compared to the baseline group: P-wave decreases from 14.3 ms (T-Base) to 9.1 ms (T-Dr2), S-wave from 14.3 ms to 11.7 ms (T-MD2, T-Dr3), and P-S time difference from 12.0 ms to 7.1 ms (T-MD3). Among uncertainty-aware T-MD series, T-MD1 showed degraded S-wave performance compared to the baseline group, while T-MD2/3 exhibited significant improvements. Notably, T-

MD3 achieves about 40% reduction in the SD of P-S time difference versus T-Base, attaining accuracy comparable to the multi-trace cross-correlation method (Tan and He, 2016). These results confirm that the proposed MD-Net framework successfully balances double objectives: maintaining phase detection and arrival picking accuracy meeting monitoring requirements while extracting crucial uncertainty metrics. The achieved precision enhancement, particularly in critical P-S time difference measurements, ensures reliable data for subsequent processing workflow.

**Fig. 10.** Phase detection and arrival picking results of different groups (double-phase example 3): (a) three component single-trace seismogram, (b) T-Base, (c) T-Dr1, (d) T-MD1, (e) T-Dr2, (f) T-MD2, (g) T-Dr3, and (h) T-MD3.



**Fig. 11.** Uncertainty quantification of T-MD series (double-phase example 3): (a) three component single-trace seismogram, (b) T-MD1, (c) T-MD2, and (d) T-MD3.

**Table 3**  
Statistics on the detection results of double-phase samples.

Group	T-Base	T-Dr1	T-MD1	T-Dr2	T-MD2	T-Dr3	T-MD3
Number of correctly detected traces	1244	1259	1253	1275	1268	1270	1267
Percentage, %	97.57	98.75	98.27	100.00	99.45	99.61	99.37

### 3.3. Location optimization evaluation

The location optimization evaluation focuses on fracturing stages 8–10 in Fig. 4(c), containing 85 double-phase microseismic

event samples. These events undergo location processing through the workflow shown in Fig. 3, combining automated data-driven phase detection and arrival picking based MD-Net and manual quality control, denoted as D-MD mode. Given the limited inter-

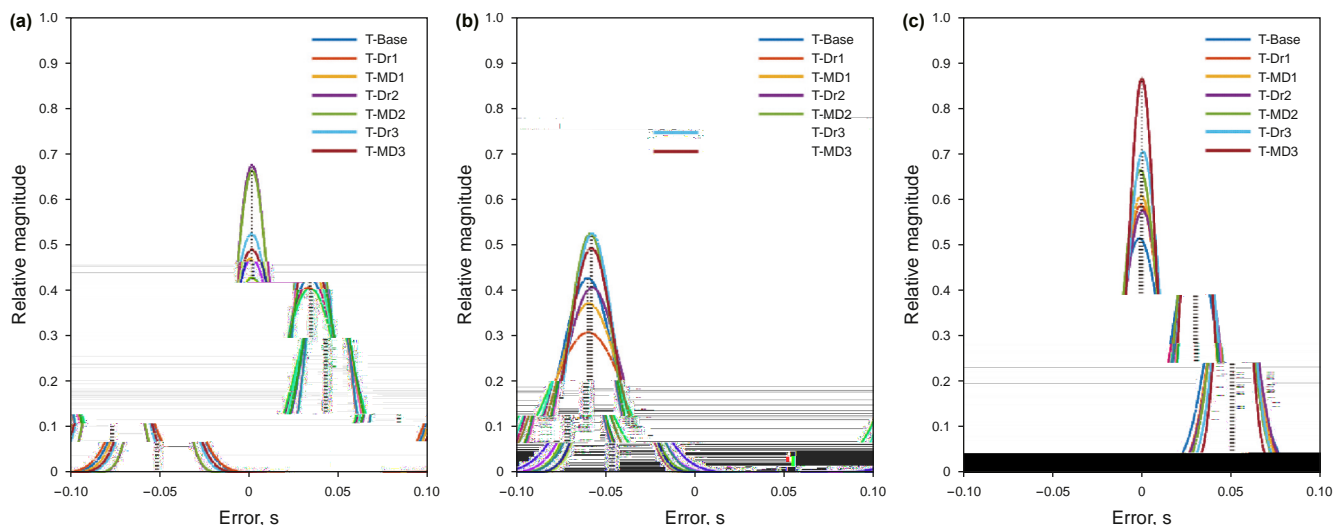


Fig. 12. Gaussian distributions relative to manual benchmarks of phase arrival picking errors of double-phase samples: (a) P-wave, (b) S-wave, and (c) P-S time difference.

well distances and small surface elevation variations, we implement a horizontal laminar velocity model derived from fracturing well sonic logging data, calibrated using perforation events from stages 8–10 (Fig. 13). Field implementation addresses uncertainty variability across dropout rates through parallel model operations: Three MD-Net models with dropout rates 0.1, 0.2, and 0.3 generate simultaneous predictions. Automated quality control triggers manual calibration when the uncertainty estimation of any model's probability distribution exceeds 0.2 (empirically determined threshold).

To verify the workflow performance of our proposed D-MD mode, three control groups are set up to carry out identical source

Table 4  
Statistics of phase arrival picking errors of double phase samples (mean: ME; standard deviation: SD).

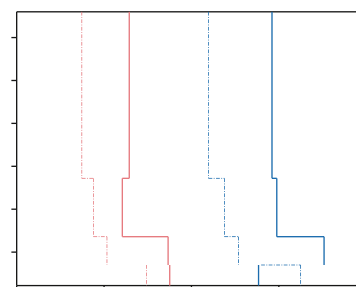
Group	P-wave		S-wave		P-S time difference	
	ME, ms	SD, ms	ME, ms	SD, ms	ME, ms	SD, ms
T-Base	0.5	14.3	1.1	14.3	1.7	12.0
T-Dr1	0.1	15.1	0.8	19.9	0.5	10.4
T-MD1	0.2	13.1	0.5	16.5	0.7	10.1
T-Dr2	1.4	9.1	1.4	15.0	0	10.7
T-MD2	1.4	9.3	0.3	11.7	1.1	9.3
T-Dr3	1.1	11.7	1.4	11.7	0.3	8.7
T-MD3	1.2	12.5	1.0	12.4	0.3	7.1

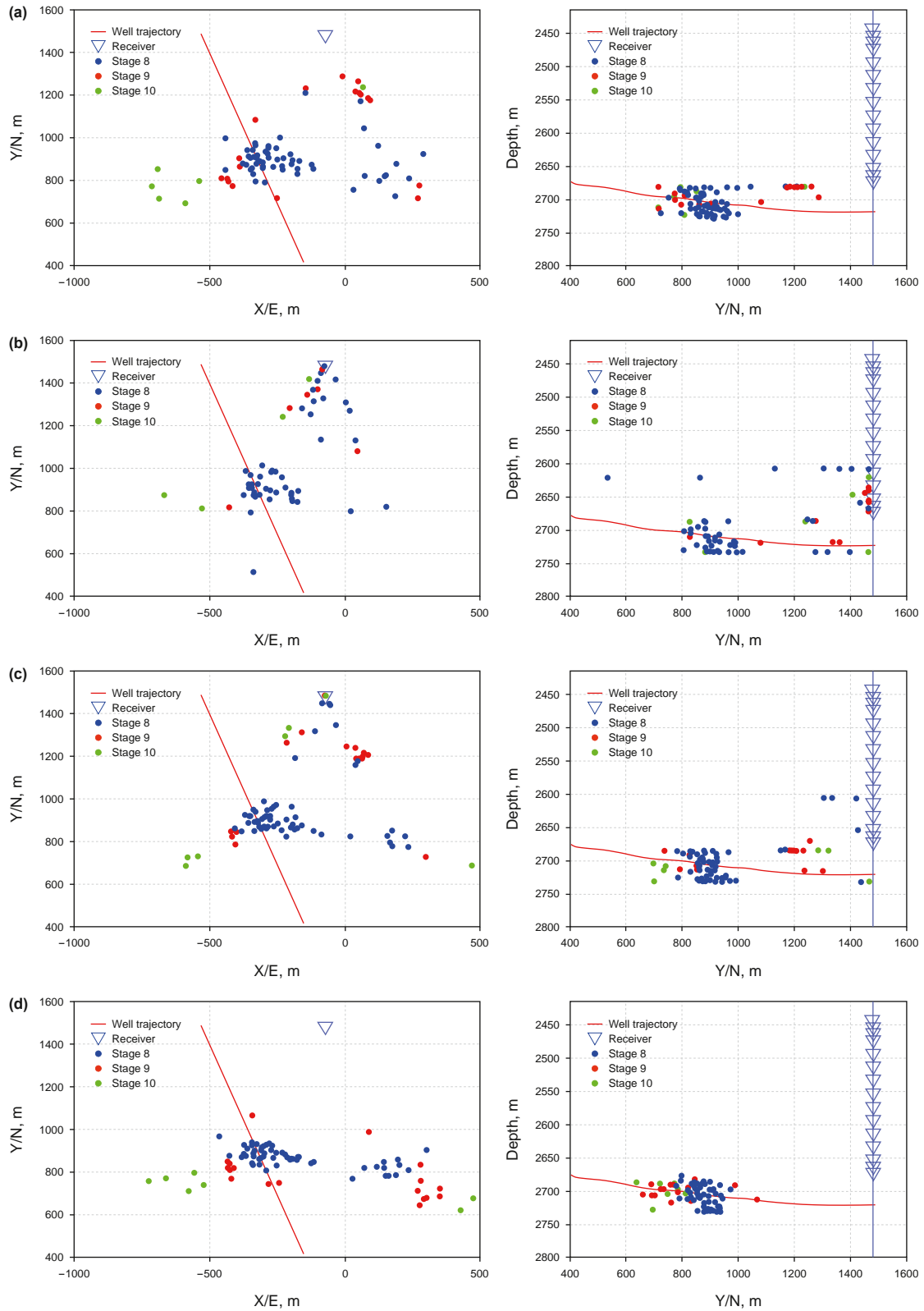


Fig. 13. Velocity model calibration using perforation events: (a) fracturing stage 8; (b) fracturing stage 9; and (c) fracturing stage 10. Blue and red lines denote velocity models of P- and S-wave respectively, with dashed lines representing initial models and solid lines indicating calibrated models.

location work without manual quality control after obtaining picking results from one model-driven (cross-correlation method) (Tan and He, 2016), two data-driven method (ST-Net (Wu et al., 2019) and MT-Net (Ma et al., 2020; Zhang et al., 2021, 2022)), denoted as M-CC mode, D-ST mode, D-MT mode, respectively. Among them, the M-CC mode, which uses waveform similarity characteristics by calculating correlation functions among multi-trace recordings, is currently recognized as one of the best strategies to improve the picking accuracy. The D-ST mode is based on the 1D U-Net structure with single-trace waveforms as input, and extracting the waveform characteristics using multilevel feature fusion of the U-Net. The D-MT mode is still based on the 2D U-Net structure, and uses multi-trace waveforms as a whole to adaptively learn in “intra-trace” and “inter-trace” dimensions meanwhile. The rest of settings for the D-ST and D-MT modes are identical to those of the D-MD except that the MC-Dropout strategy is not used. All modes share identical velocity models and azimuthal constraints, with source locations determined through grid search inversion.

Fig. 14(a) presents the source location results obtained through the conventional M-CC mode. Spatial analysis reveals distinct clustering patterns: Microseismic events of stage 8 predominantly concentrate near wellbores, events of stage 9 show partial location adjacent to monitoring wells, while events of stage 10 demonstrate sparse distribution with limited spatial coherence. The D-ST mode (Fig. 14(b)) outputs exhibit significant location anomalies, with parts of events erroneously localized beyond target fracturing





**Fig. 14.** Spatial distribution of microseismic sources using different modes: (a) M-CC, (b) D-ST, (c) D-MT, (d) D-MD. Hypocenters from fracturing stages 8, 9, and 10 are denoted by blue, red, and green scatter points respectively.

zones (exceeding grid search boundaries), likely attributed to errors in P-S time difference measurements of some traces of those events. Although the D-MT mode (Fig. 14(c)) demonstrates

improved clustering along fracture expanding direction, minority of events remain mislocated near monitoring wellbore, resulting in loss of microseismic data for fracture interpretation. The

proposed D-MD mode (Fig. 14(d)) achieves optimal performance with all of events concentrated along the primary fracture expanding direction near horizontal wellbore. Compared to other modes, D-MD mode avoids location errors while maintaining geomechanically consistent spatial distributions. This precision enhancement confirms the effectiveness of our proposed workflow in meeting practical requirements for real-time fracturing interpretation of HF scenarios.

#### 4. Discussion

The integration of Monte Carlo Dropout into the MD-Net architecture introduces a paradigm shift in uncertainty quantification for seismic phase detection and arrival picking. By approximating Bayesian inference through stochastic forward sampling, the model not only suppresses noise-induced picking errors but also generates confidence intervals that guide manual calibration. The D-MD series maintains computational efficiency comparable to other experimental groups during training, with only a marginal increase in prediction cost due to multiple sampling. For field microseismic monitoring, this added latency (~0.1 s/event) is negligible compared to operational timescales and does not hinder practical deployment. The framework thus balances uncertainty quantification with minimal computational overhead in field applications.

Furthermore, the hybrid workflow underscores the synergy between data-driven automation and physics-based constraints. Unlike pure data-driven approaches, which may produce geologically implausible outliers, the incorporation of traveltimes-based inversion ensures adherence to wave propagation principles, bridging the gap between black-box predictions and interpretable geophysical models. The current study is designed for extensibility to diverse lithologies and noise profiles. Zhang et al. (2022) demonstrated that transfer learning enables effective domain adaptation across distinct fracturing projects and geological regions. By integrating similar TL strategies, MD-Net can be rapidly fine-tuned for new scenarios using minimal target data, significantly reducing deployment barriers.

Despite its strengths, the framework exhibits three primary limitations:

- (1) The dropout rate significantly influences uncertainty estimation robustness. Excessively high rates may oversmooth phase characteristic, while low rates inadequately capture epistemic uncertainty. Future implementations could integrate Bayesian optimization to dynamically adjust dropout rates based on real-time SNR assessments.
- (2) The current workflow assumes laminar velocity models calibrated via perforation events. In complex reservoirs, this simplification may introduce location biases. Coupling MD-Net integrating distributed acoustic sensing (DAS) data could address such limitations (Tegtow et al., 2025).
- (3) The microseismic events processed in field case are events with clear phase arrivals, and the picking error is small. Therefore, the location results using traditional grid search algorithm can basically meet

## Declaration of competing interest

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests: Chuan He reports financial support was provided by DeepEarth Probe and Mineral Resources Exploration - National Science and Technology Major Project (Grant No. 2024ZD1002503). If there are other authors, they declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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## References

- Akram, J., Eaton, D.W., 2016. A review and appraisal of arrival-time picking methods for downhole microseismic data. *Geophysics* 81 (2), KS71–KS91. <https://doi.org/10.1190/geo2014-0500.1>.
- Anikiev, D., Birnie, C., Waheed, U., et al., 2023. Machine learning in microseismic monitoring. *Earth Sci. Rev.* 239. <https://doi.org/10.1016/j.earscrev.2023.104371>.
- Bao, X., Eaton, D.W., 2016. Fault activation by hydraulic fracturing in Western Canada. *Science* 354. <https://doi.org/10.1126/science.aag2583>.
- Bergen, K.J., Beroza, G.C., 2019. Earthquake fingerprints: extracting waveform features for similarity-based earthquake detection. *Pure Appl. Geophys.* 176, 1037–1059. <https://doi.org/10.1007/s00024-018-1995-6>.
- Bergen, K.J., Johnson, P.A., de Hoop, M.V., et al., 2019. Machine learning for data-driven discovery in solid Earth geoscience. *Science* 363. <https://doi.org/10.1126/science.aau0323>.
- Chen, Y., Saad, M., Savvaidis, O., Alexandros, et al., 2024. Deep learning for P-wave first-motion polarity determination and its application in focal mechanism inversion. *IEEE Trans. Geosci. Rem. Sens.* 62. <https://doi.org/10.1109/TGRS.2024.3407060>.
- Gal, Y., Ghahramani, Z., 2016. Dropout as a Bayesian appro

Zheng, M., Li, J., Wu, X., et al., 2018. China's conventional and unconventional natural gas resources: potential and exploration targets. *Journal of Natural Gas Geoscience* 3 (6), 295–309. <https://doi.org/10.1016/j.jnggs.2018.11.007>.

Zhou, Y., Ding, H., Ghost, A., et al., 2025. AI-PAL: self-supervised AI phase picking via rule-based algorithm for generalized earthquake detection. *J. Geophys. Res. Solid Earth* 130 (4). <https://doi.org/10.1029/2025JB031294>.

Zhou, Y., Yue, H., Kong, Q., et al., 2019. Hybrid event detection DZj pDZípS8SSSSDZvp.8.)SxZS